

MARINE RESERVOIR VARIATION IN THE BISMARCK REGION: AN EVALUATION OF SPATIAL AND TEMPORAL CHANGE IN ΔR AND R OVER THE LAST 3000 YEARS

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ABSTRACT. Interactions between islands, ocean currents, and winds cause large-scale eddies and upwelling in the lee of islands that can result in spatial variation in the marine radiocarbon reservoir. For waters around New Ireland and the Bismarck Sea, ΔR values ranging from 365 to -320 ^{14}C yr have been reported (Kirch 2001; Petchey et al. 2004). Petchey et al. (2004) proposed that some of this variation was caused by seasonal reversals in the South Equatorial Current and North Equatorial Counter Current system, combined with Ekman upwelling from the Equator. McGregor et al. (2008) suggested additional complexity within this region caused by a change in the reservoir value over time in response to changing climatic conditions. We present a series of 14 new and extant published ΔR and R values on historic shells, combined with 8 values from archaeological terrestrial/marine pairs and U-Th dated coral, that support observations of localized variability caused by a complex interplay between seasonal currents, riverine input, and ocean eddies. On the basis of these values and oceanographic data, we divide the Bismarck Sea surface marine ^{14}C reservoir into 6 tentative subregions. In particular, our results support significant variation within channels at the southwest and southeast ends of New Britain and towards the equatorial boundary of the sea. Our results indicate that within the Bismarck Sea geographical variation appears to be more extreme than temporal over the last 3000 yr.

INTRODUCTION

For archaeologists and Quaternary researchers working in the Bismarck region, it is critical that reliable calibrated ages can be obtained from marine shell because, like most archaeological sites in the Pacific, charcoal is often highly degraded, unidentifiable, and/or scarce (see Ambrose 1988; Specht and Gosden 1997; Kirch 2001:201–2; Summerhayes 2007; Specht 2009). The global average marine reservoir age ($R_g(t)$) (i.e. the difference between the atmospheric ^{14}C content and that of the surface ocean) of surface water, and shells by association, is around 400 ^{14}C yr (Stuiver et al. 1986). By convention, this offset is corrected for when a marine radiocarbon date is calibrated using the appropriate marine curve (e.g. Marine09; Reimer et al. 2009). Regional values that deviate from $R_g(t)$ are often expressed as ΔR (the difference between the modeled ^{14}C age of surface water and the known ^{14}C age of surface water at that time). The reservoir age correction value (R) is also used by some researchers (R is defined as the offset between ^{14}C age of the sample and the contemporaneous regional atmosphere at any given time). However, because the oceanic response to atmospheric ^{14}C forcing differs from the atmospheric signal, R will vary over time. Consequently, R cannot be used to distinguish regional differences of ± 100 yr unless carbon in the samples is contemporaneous (Stuiver et al. 1998:1131). To assess the geographical variation in the marine reservoir, it is necessary therefore to use ΔR because the regional ocean and the world ocean response to atmospheric forcing are assumed to be constant and cancel out any temporal change in the marine reservoir ^{14}C .

ΔR can be calculated from marine samples collected from known locations prior to AD 1950, from contemporaneous terrestrial/marine samples such as those found in archaeological deposits (e.g. Reimer et al. 2002; Ulm 2002; Petchey et al. 2005; Petchey and Clark 2011), or paired U-Th ^{14}C dates from corals (Edwards et al. 1993; McGregor et al. 2008). Although the database of ΔR values for the Pacific is growing (<http://calib.quub.ac.uk/marine/>), the interpretation and application of these ΔR values to marine shell calibrations is of limited use without an evaluation of regional ocean

circulation patterns, local marine conditions, as well as the habitat and dietary preferences of the shellfish dated (Tanaka et al. 1986; Hogg et al. 1998; Petchey and Clark 2011). In the wake of this increased interpretive complexity, shell ^{14}C dates have obtained a reputation for unreliability. Some recent reviews of ^{14}C data have ignored shell altogether (e.g. Wilmschurst et al. 2010; Rieth et al. 2011), while others have overlooked the inherent regional variability by the application of a single ΔR (e.g. Specht and Summerhayes 2007:55; White 2007:5; Summerhayes 2010).

Recent research has indicated that considerable ΔR variation is possible not just in estuarine or lagoon environments where terrestrial ^{14}C input is likely (Ulm et al. 2009; Petchey and Clark 2011), but also where ocean currents meet continental landmasses, opposing currents, or island chains (Petchey et al. 2008). The Bismarck Sea has previously been identified as problematic in this respect (Petchey et al. 2004, 2005; McGregor et al. 2008). In this region, water derived from the southern branch of the South Equatorial Current (SEC) flows around the southeast tip of Papua New Guinea (PNG) and enters the Bismarck Sea through the Vitiaz Strait—between New Britain and PNG—as the New Guinea Coastal Current (NGCC) (Figure 1a,b). Further north, the flow of the more southerly branch of the SEC is broken by the Solomon Islands before entering the Bismarck Sea either through the St. George's Channel, located between New Britain and New Ireland, or the Ysabel Channel between New Ireland and Manus. The dominant northwest flow (NGCC; June/July) is reversed during the winter (January/February) monsoon and the NGCC flows back through the Vitiaz and St. George's channels (Kuroda 2000).

A significant change in the marine ^{14}C reservoir has also been identified in the period 7220–5850 BP ($\Delta\text{R} = -105 \pm 110$ to -265 ± 35 ^{14}C yr compared with 70 ± 60 ^{14}C yr¹ for a modern coral from Muschu Island). McGregor et al. (2008) attributed this to strengthening tradewinds and a shift in the Intertropical Convergence Zone, resulting in a greater influx of well-equilibrated subtropical water. McGregor et al. (2008:220–1, Table 8) also recorded an abnormal shift in ΔR to -135 ± 35 ^{14}C yr around 2065 yr ago. Recently, Yu et al. (2010) noted a shift in ΔR for corals from the South China Sea dating to around 2800 to 2500 yr ago, but of opposing magnitude (180 ± 20 ^{14}C yr and 103 ± 20 ^{14}C yr) to that recorded by McGregor et al. (2008). This variability remains of potential concern to archaeologists investigating human expansion at this time.

In this paper, we present results of research aimed to investigate further variation in the marine reservoir, geographically and over time in the Bismarck region, using terrestrial and marine materials from archaeological sites, U-Th dated corals, and shells from known locations collected in the 19th and early 20th centuries.

METHODS

Six pre-AD 1950 shell ^{14}C values and 1 pre-AD 1955 shell ^{14}C value are presented here for the first time. These were obtained from the Auckland War Memorial Museum in New Zealand and the Australian Museum in Sydney. An additional 8 values have been obtained from published works (Chappell and Polach 1976; Petchey et al. 2004; McGregor et al. 2008; Burr et al. 2009; Ambrose et al., in press) (Table 1). These historic values are compared to 4 archaeological shell/charcoal pairs from Watom Island (Green and Anson 2000; Petchey et al. 2005); Mussau (Kirch 2001) and Sasi in the Lou Islands (Ambrose 1988) (Table 2); and 4 paired U-Th/ ^{14}C ages on coral (Edwards et al. 1993; McGregor et al. 2008) (Table 3). The location of each site and shell/coral collection point is shown in Figure 1a.

¹These values are taken directly from McGregor et al. (2008) and differ from values presented in Table 1, which have been recalculated according to the methodology outlined here.

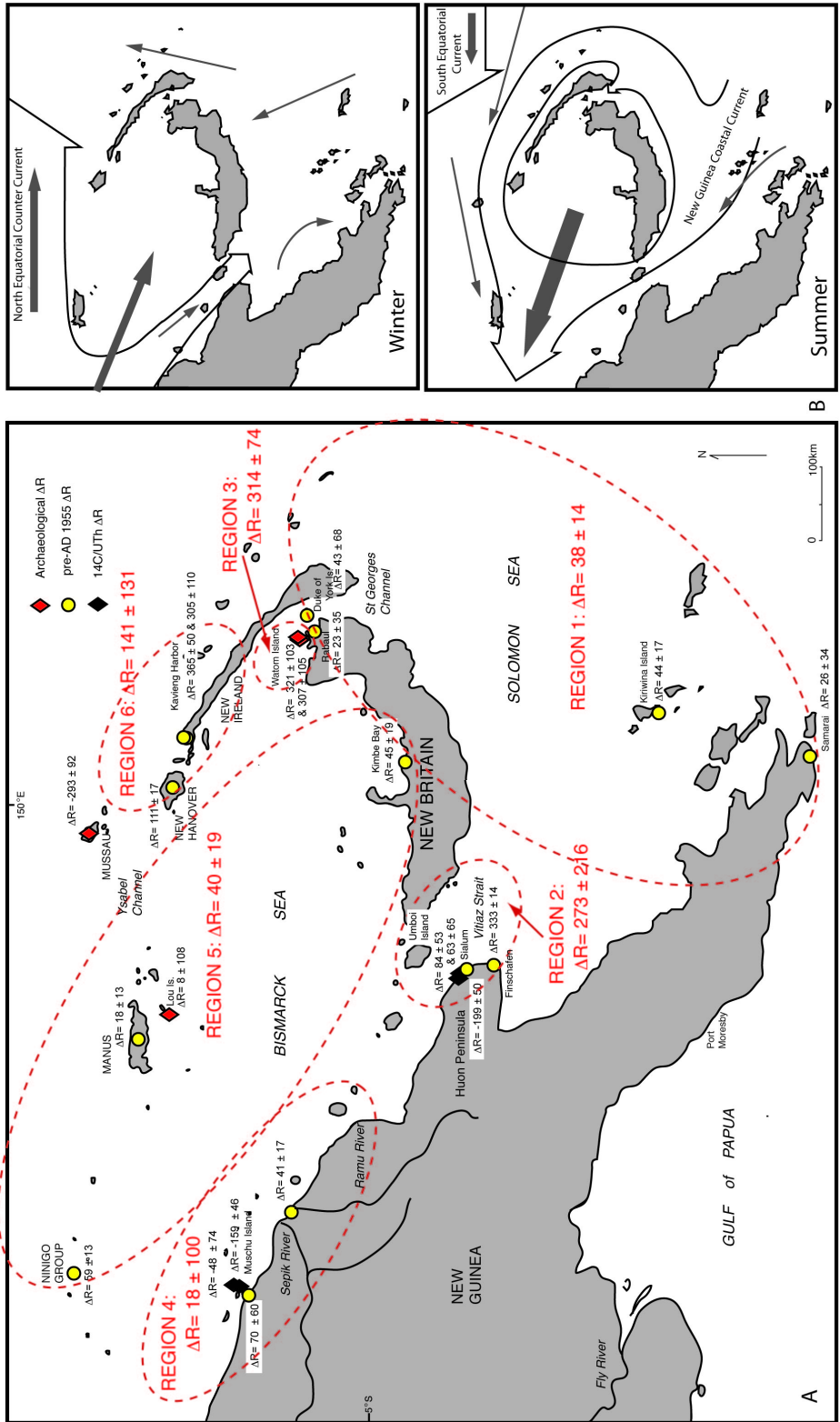


Figure 1 A) Map showing location of marine samples and associated ΔR values mentioned in the text. The map is divided into 6 ΔR subregions based on oceanographic observations. B) Major current patterns during 1999/2000 for summer (June/July) and winter (Jan/Feb) [Current data from AIMS (1998)].

When selecting a suitable sample for ΔR work, it is essential that the age of shellfish death is known (Stuiver and Braziunas 1993). For museum specimens, this information should be clearly documented. For archaeological ΔR , this requirement is determined by dating charcoal from short-lived plants in contexts that are contemporaneous with the shell selected. Lastly, the shell must be devoid of possible heirloom effects (e.g. food shells are usually discarded at the time of collection) or inbuilt age. Unfortunately, although a large number of archaeological dates are available in this region, few pairs conform to these strict guidelines. The most common problem involves the use of unidentified charcoals that could have inbuilt age (i.e. the growth age of the tree) (see Anderson et al. 2001; Allen and Wallace 2007) or storage age (i.e. where a period of time has elapsed from death to use by people; McFadgen 1982). Fortunately, stored wood is unlikely in the humid tropics where wood decays rapidly (Swift et al. 1979; Kirch 2001:202). This is, however, not the case for shell artifacts. Many sites also have a long occupation history, and locating suitable samples from undisturbed contexts can be problematic (see Petchey et al. 2009). The impact of these factors on ΔR varies. Inbuilt age in charcoal will reduce the difference between shell and charcoal ^{14}C determinations and the ΔR value will be smaller and inaccurate. Any storage age in the shell samples will result in an incorrect and larger ΔR value (more positive), assuming the associated charcoal pair is short-lived.

Shellfish Habitat and Dietary Considerations

Dating shellfish and assigning a ΔR can be complex because the ^{14}C content of an animal is not always in equilibrium with the ocean water in which they live. Approximately 50% of carbon in the shells of marine invertebrates is derived from metabolic sources (Tanaka et al. 1986). Because suspension-feeding shellfish predominantly consume suspended phytoplankton² and dissolved inorganic carbon (DIC) from seawater, it is generally assumed, therefore, that they reflect surface ocean reservoir conditions (Hogg et al. 1998). It is, however, possible that carbon from sources other than ocean DIC can become incorporated in the shells (cf. Keith et al. 1964; Dye 1994; Petchey and Clark 2011). Therefore, it is important that all ΔR results are evaluated with possible variations in diet, habitat, and local oceanic conditions kept in mind (Bondevik and Gulliksen *in* Mangerud et al. 2006: 3241).

The majority of shellfish given in Tables 1, 2, and 3 are suspension feeders (*Anadara* sp., *Begonia* sp., *Comptopallium* sp., *Barbatia* sp., *Meretrix* sp., *Pinctada* sp., *Tridacna* sp., and *Hyotissa* sp.). In addition to suspension feeding, the giant clams (*Tridacna* sp.) also obtain energy from carbohydrates obtained from photosynthesis via a symbiotic relationship with zooxanthellae (single-celled algae) (Beesley et al. 1998:333–5). *Trochus* sp. and *Nerita* sp. are both algal grazing herbivores (Beesley et al. 1998:683, 697), and *Nerita* has previously been identified as having problematic ^{14}C values when recovered from areas of limestone bedrock due to the ingestion of terrestrial sediment while grazing (Anderson et al. 2001). *Nassarius* sp. are carnivorous and occupy a wide range of marine environments (Beesley et al. 1998:831, 852–3). Isotopic studies (Cook et al. 2004:882; Petchey and Clark 2011) indicate that their values will reflect their position as a secondary consumer.

With the exception of the giant clams that, depending on species, may live for 100 yr (Hamel and Mercier 1996), the shellfish listed in Tables 1 and 3 all live for less than 50 yr. Because the ΔR results for the giant clam specimens from archaeological deposits (Table 2) come from published sources, we cannot determine if the dates are of juvenile or adult specimens; however, excessive age would result in ΔR values that are very large, which is not the case in this instance (see below).

²Phytoplankton are primary producers in the marine environment, deriving their carbon from DIC.

Table 1 New and extant radiocarbon data for known-age, pre-AD 1955 shells and corals. See Table 3 bottom for footnotes for Tables 1–3.

Location	Sample material ^a	Historical age (AD)	Marine		ΔR (yr)	R [Rs(t) marine– ¹⁴ C atmos]	Identifier ^c	Collector/ Reference
			¹⁴ C age & error (BP) [Rs(t)]	modeled age ^b [Rg(t)]				
NEW ΔR VALUES								
Kalolo Cape, Hoskins Peninsula, Kimbe Bay	<i>Anadara</i> sp. (SF)	1944	508 ± 19	463 ± 23	45 ± 19	323 ± 21	Wk-19686 [AM:C449012] ^d	H Consett Davies
Finschhafen, Huon Peninsula	<i>Barbatia</i> sp. (SF)	1944	796 ± 14	463 ± 23	333 ± 14	611± 16	Wk-20352 [AM:C449015] ^d	H Consett Davies
Ninego Group	<i>Beguinia semiorbiculata</i> (SF)	~1919	508 ± 13	449 ± 23	59 ± 13	384 ± 15	Wk-20350 [AM:C045938] ^d	Capt W Burrows
New Hanover	<i>Pinctada maculata</i> (SF)	~1919	560 ± 17	449 ± 23	111 ± 17	436 ± 18	Wk-21067 [AM:C045975] ^d	Capt W Burrows
Manus Island	<i>Comptopallium radula</i> (SF)	1954	487 ± 13	469 ± 24	18 ± 13	~288 ± 16 ^v	Wk-20351 [AM:C132142] ^d	L Woolcott
Near mouth of Ramu River	<i>Meretrix</i> sp. (SF)	1925–26	492 ± 17	451 ± 23	41 ± 17	360 ± 18	Wk-21068 [Ak-] ^d	R W Gosset
EXTANT ΔR VALUES								
Samarai	<i>Barbatia</i> sp. (SF)	1938	485 ± 34	459 ± 23	26 ± 34	320 ± 35	Wk-26249 [NMNZ 212226] ^d	W R B Oliver (Ambrose et al., in press)
Kiriwina Island, Trobriand Group	<i>Pinctada imbricata</i> (SF)	1944	507 ± 17	463 ± 23	44 ± 17	322 ± 19	Wk-21066 [AM:C119161] ^d	H Consett Davies (Ambrose et al., in press)
Huon Peninsula, west of Stialum	Historic (coral) <i>Flavia</i> (~5 rings)	1955	270 ± 50	469 ± 24	–199 ± 50	71 ± 51 ^e	ANU-114	Chappell and Polach (1976:237)
Rabaul Harbor	Coral: <i>Porites</i> sp.	1945–1950	490 ± 35	467 ± 24	23 ± 35	297 ± 36	AA-series	Burr et al. (2009:327)
Duke of York Is	<i>Nassarius camelus</i> (C)	~1905 ^f	492 ± 68	449 ± 23	43 ± 68	408 ± 68	Wk-9219	Petchey et al. (2004)
Kavieng Harbor ^g	<i>Nerita plicata</i> (H)	1931	820 ± 50	455 ± 23	365 ± 50	668 ± 51	Wk-8377	Petchey et al. (2004)
Kavieng Harbor	<i>Barbatia foliata</i> (SF)	1931	760 ± 110	455 ± 23	305 ± 110	608 ± 110	Wk-8379	Petchey et al. (2004)
Muschu Island, Cape Saum	Coral: <i>Porites lutea</i>	1912	520 ± 60	448 ± 23	72 ± 60	419 ± 60	MS01	McGregor et al. (2008:217)

Table 2 Radiocarbon data for paired charcoal/shell samples from archaeological sites.

Location	Sample material ^a	¹⁴ C age & error (BP) [Rs(t)]	Pooled CRA error (BP)	Marine modeled age ^b (Rg(t))	ΔR (yr) Rs(t)–Rg(t)	R [Rs(t) marine– ¹⁴ C atmos]	Identifier ^c	Reference
Site SASI – Lou Islands, Admiralty Islands								
Site 9(4)	Charcoal (unidentified)	2190 ± 100	2105 ± 49	2442 ± 61	8 ± 108	345 ± 103	ANU-2155	Ambrose (1988:489)
Sasi soil	Charcoal (unidentified)	2070 ± 80	[$\chi^2_{3,0.05} = 0.97 < 7.81$]				ANU-2014	
Sasi soil	Charcoal (unidentified)	2090 ± 100					ANU-5398	
Sq E, 0–10 cm	Charcoal (twig + some larger fragments)	2080 ± 130					Wk-8544	Unpublished data
Sasi soil	<i>Tridacna</i> sp. (SF/Ph)	2480 ± 90	2450 ± 90	—			ANU-5399	Ambrose (1988:489)
Sasi soil	Shell (unidentified)	2300 ± 90	[$\chi^2_{2,0.05} = 3.33 < 5.99$]				ANU-4981	
Sq E, 0–10 cm	<i>Anadara</i> sp. (SF)	2480 ± 45					Wk-8545	Unpublished data
Site ECA – Mussau^h								
ECA, Zone C3 of Area B	<i>Intsia bijuga</i> outer rings	2950 ± 80	Mean = 2940 ± 57 [$\chi^2_{1,0.05} = 0.03 < 3.84$]	3273 ± 72	–293 ± 92	40 ± 81	ANU-5790	Kirch (2001:226)
ECA, Zone C3 of Area B	<i>Intsia bijuga</i> outer rings	2930 ± 80					ANU-5791	
Unit W200N151, Level 11, Zone C3	<i>Tridacna gigas</i> (SF/Ph)	3010 ± 80	2980 ± 57 [$\chi^2_{1,0.05} = 0.28 < 3.84$]	—			ANU-5081	Kirch (2001:201, 225–6)
Unit W201N149, Level 12, Zone C3	<i>Hyotissa hyotis</i> (SF)	2950 ± 80					ANU-5082	
Site SAC – Watom Island								
Sq F15, spit 1	Coconut charcoal	1730 ± 60	—	2069 ± 65	321 ± 103	660 ± 100	Wk-7371	Petchey et al. (2005); Green and Anson (2000)
Sq G14 (base of layer)	<i>Tridacna maxima</i> (SF/Ph)	2390 ± 80	—				ANU-5330	
Sq G13, spit 2	Coconut charcoal	2860 ± 60	—	3183 ± 68	307 ± 105	630 ± 100	Wk-7370	
Sq G10, 1.85 m below surface	<i>Trochus niloticus</i> (H)	3490 ± 80	—				ANU-5339	

Table 3 Radiocarbon data for paired U-Th/coral samples.

Location	Sample material ^a	Calendar age (BP)	¹⁴ C age & error (BP)		Pooled CRA error (BP)	Marine modeled age ^b		ΔR (yr)		R	Identifier ^c	Reference
			[Rs(t)]	[Rg(t)]		[Rs(t)]	[Rg(t)]	[Rs(t)]	[Rg(t)]			
Muschu Island – Rebiew Bay	Coral	1940 ± 50 (U-Th)	2280 ± 60	—	—	2328 ± 44	—48 ± 74	289 ± 72	—	ANU-11056	McGregor et al. (2008:219)	
	Coral	2180 ± 50 (U-Th)	2370 ± 35	—	—	2529 ± 30	−159 ± 46	65 ± 62	—	[FM09] ^d OZI-426 [FM09] ^d		
Huon Pennin- sula (5 km south of Sialum)	Coral	1808 ± 9 ⁱ	1793 ± 20 [$\chi^2_{1:0.05} =$ 5.19<3.84]	2325 ± 40 [$\chi^2_{1:0.05} =$ 2.61<3.84]	2268 ± 48	2184 ± 22	84 ± 53	432 ± 49	—	AA-series [HOB-EE-5a] ^d	Edwards et al. (1993:963); G Burr, pers. comm., Jan 2011	
		1779 ± 9	2252 ± 21	—	—	—	—	—	—	WHAMS-977 [HOB-EE-5b] ^d	E Druffel, pers. comm., Feb 2011	
	Coral	2350 ± 11	2760 ± 60	—	—	2697 ± 25	63 ± 65	406 ± 93	—	AA-series [KAI-EE-1a] ^d	Edwards et al. (1993:963); G Burr, pers. comm., Jan 2011	

^aDiet preferences (in brackets): SF = suspension-feeder; C = carnivore; H = herbivore; SF/Ph = mixed suspension feeding and zooxanthellae photosynthesis.

^bBecause mollusk shells are built up over their entire lifespan, the margins of the shell will be younger than the hinge. For new ΔR , values we have sampled no more than 5 “circuli.” Therefore, when calculating the marine modeled age for these shells we have assumed that the carbon was fixed into the shells close to the year of collection. Where necessary, we have interpolated between the 5-yr increments in the Marine09 ages.

^cLab prefixes: Wk = Waikato Radiocarbon Dating Laboratory; CAMS = Lawrence Livermore National Laboratories; AA = University of Arizona; OZI = Australian Nuclear Science and Technology Organisation; WHAMS = Woods Hole Oceanographic Institution; ANU = Australian National University.

^dMuseum acquisition number: AK = Auckland War Memorial Museum, New Zealand; AM:C = Australian Museum, Sydney, Australia; NMNZ = National Museum of New Zealand – Te Papa.

^eWe have assumed the ¹⁴C content of post-AD 1950 samples (Wk-20351 and ANU-114) are equivalent to the AD 1950 atmosphere (= 199 ± 9 ¹⁴C yr).

^fWk-9219 was collected prior to AD 1910 and we have assumed an AD 1905 date.

^gOZB-768 to OZB-773 from Kavieng (published in Petchey et al. 2004) have been re-inspected and are no longer considered to have been live collected.

^hThree values from ECB have been excluded from Table 2. This includes charcoal date Beta-20453 (3200 ± 70 BP) because it was not on identified short-lived material, and paired shell dates (*Hyotissa hyotis*) ANU-5086 (3120 ± 80 BP) and ANU-5087 (3150 ± 80 BP). If the associated charcoal can be confirmed to be short-lived material, these values would result in a ΔR of −406 ± 81 ¹⁴C yr and a R of −65 ± 90 ¹⁴C yr.

ⁱ¹⁴C result differs from that in previous publications because Edwards et al. (1993:963) corrected for R by subtracting 407 ± 52 ¹⁴C yr. The value given in Table 3 also includes additional ¹⁴C results obtained post-AD 1993.

Both R and ΔR values are given in Tables 1–3. For each of the ^{14}C results, the ΔR for a specific location “(s)” was calculated using the formula $R_s(t) - R_g(t) = \Delta R(s)$, where $\Delta R(s)$ is the difference between the actual ^{14}C activity of the surface ocean at a particular location [$R_s(t)$] at that time and the global average [$R_g(t)$] (Marine09: Reimer et al. 2009) (Stuiver et al. 1986). For historic shell and coral samples, the ΔR standard error is the 1σ estimate of uncertainty in the conventional ^{14}C age of the shell sample; this avoids duplication of the uncertainty in the Marine09 curve when the ΔR is used for calibration (P Reimer, personal communication, 2007). R is the difference between the ^{14}C age of the shell/coral sample [$R_s(t)$] and the atmospheric ^{14}C age (Table 1).

For archaeological pairs (Table 2), an estimate of the atmospheric calibration curve error (IntCal09: Reimer et al. 2009) over the 1σ span of the ^{14}C age was used to derive the calculated marine modeled age [$R_g(t)$] error, whereby, atmospheric age $\sigma = (\sigma^{14}\text{C age}^2 + \text{average of calibration curve error}^2)$. The ΔR standard error is calculated by the formula $\Delta R\sigma = (\sigma R_g(t)^2 + \sigma R_s(t)^2)$. Similarly, the U-Th coral ages (Table 3) were converted to atmospheric ^{14}C ages using the IntCal09 curve whereby the equivalent atmospheric ^{14}C age is the median of the oldest and youngest ^{14}C ages across the U-Th (1σ) age range.

All R values have been calculated using the IntCal09 data sets (Reimer et al. 2009). The use of the Northern Hemisphere data set is based on the recommendations of McCormac et al. (2004:1088) and Petchey et al. (2009) whereby the ^{14}C boundary between the atmosphere of the Southern and Northern hemispheres lies along the thermal equator, commonly called the Intertropical Convergence Zone. This varies from the procedure used by Burr et al. (2009) for this region whereby reservoir ages were calculated by subtracting Southern Hemisphere atmospheric data derived from McCormac et al. (2004) and Hua et al. (2000, 2003).

Even when samples are carefully selected according to the prerequisites listed above, there are a number of uncertainties with the ^{14}C data because of the postulated time of carbon uptake before collection, the influence of diet and habitat, and short-term fluctuation in the water masses. To estimate the amount of uncertainty that needs to be added to the ΔR value by the nonuniform ^{14}C content of the shellfish, we have calculated the weighted mean for each group (Table 1) using the chi squared (χ^2) test to evaluate the internal variability in ΔR values. Then, using the methodology recommended by Bondevik and Gulliksen in Mangerud et al. (2006), if the group has additional measurement variability (as indicated if $\chi^2/(n-1)$ is >1), we have added an additional uncertainty (external variance) to the ΔR . In this instance, the uncertainty is calculated by $\sqrt{(s^2_{\Delta R \text{ pooled}} + \sigma_{\text{ext}}^2)}$, whereby the external standard deviation (σ_{ext}) is determined by subtracting the ^{14}C measurement variance from the total population variance and obtaining the square root (e.g. $\sigma_{\text{ext}} = \sqrt{(\sigma_{\text{pop}}^2 - \sigma_{\text{meas}}^2)}$) (see Mangerud et al. 2006:3241–2 for explanation). When $\chi^2/(n-1)$ is 1, the uncertainty on the individual measurements explains the variations within the group of ΔR values and the weighted mean is used.

RESULTS

A total of 22 ΔR values from the southwest tropical Pacific dating to within the last 3300 yr are presented in Tables 1–3. ΔR values range from 365 ± 50 to -293 ± 71 ^{14}C yr—a spread of around 700 ^{14}C yr (Figure 1). The greatest variation in ΔR at 1 location occurs around the Huon Peninsula where values range from 333 ± 14 to -199 ± 50 ^{14}C yr. The most positive ΔR have been recorded for New Hanover (111 ± 17 ^{14}C yr), Kavieng Harbor at the northern tip of New Ireland (365 ± 50 and 305 ± 110 ^{14}C yr), Finschafen (333 ± 14 ^{14}C yr), and for archaeological deposits on Watom Island (307 ± 105 and 321 ± 103 ^{14}C yr). Unusually low ΔR have been measured on archaeological shells from Mussau (-293 ± 92 ^{14}C yr), located just outside of the Bismarck Sea. This range of variation within

a relatively small area means it is impractical to apply a regional ΔR value as has been possible for the central South Pacific Gyre region ($\Delta R = 5.5 \pm 21$ ^{14}C yr) (Petchey et al. 2008; Table 1). Although we recognize that this region is highly complex and the variability present is large, we argue that, when local oceanographic conditions are taken into consideration, it is possible to apply valid and statistically sound subregional ΔR values that incorporate the variation observed.

Ongoing research within the wider Pacific has indicated that wide shifts in ΔR can occur on the leeward side of islands, where fast-flowing currents and eddy formation result in the upwelling of ^{14}C -depleted waters (e.g. the NEC and NECC flows that pass Hawai'i [Petchey 2009] and Palau [Petchey and Clark 2010], respectively, that exceed 25 cm s^{-1}). Seasonal current reversals further compound this affect in the narrow channels that feed into the Bismarck Sea. Chlorophyll concentration and sea surface temperature satellite data reported in Steinberg et al. (2006) indicate intermittent upwelling along the southeast coast of New Britain, Kimbe Bay, southwest Manus, Mussau, and along the Papua New Guinea coast near the Vitiaz and George's channels. Variable, and sometimes very high, ΔR values are recorded around the Vitiaz, St. George's, and Ysabel channels in keeping with these observations (Tables 1–3). These values appear to be higher in the lee of the dominant current direction; for example, in St. George's Channel the dominant flow all year round is from the Solomon Sea to the Bismarck Sea with speeds that can reach 100 cm s^{-1} (Lindstrom et al. 1990:182). Values of 307 ± 105 and 321 ± 103 ^{14}C yr (ΔR average = 314 ± 74 ^{14}C yr; $\chi^2_{1:0.05} = 0.01 < 3.84$) for Watom Island on the western side of the channel are typical of values found in upwelling zones, and are most likely caused by eddies, while those to the southeast ($\Delta R = 43 \pm 68$ ^{14}C yr for the Duke of York Islands and 23 ± 35 ^{14}C yr for Rabaul Harbor) are lower and more in keeping with values reported by Ambrose et al. (in press) from Kiriwina and Samarai islands (southeastern Papua New Guinea) ($\Delta R = 44 \pm 17$ and 26 ± 34 ^{14}C yr, respectively) and those found in the Coral and Solomon seas (Petchey et al. 2004). The combined ΔR for the Duke of York Islands, Samarai, Kiriwina Island, and Rabaul is 38 ± 14 ^{14}C yr ($\chi^2_{3:0.05} = 0.44 < 7.81$; no external variance). This is indistinguishable from the South Pacific Gyre average of 5.5 ± 21 ^{14}C yr.³

The pattern for Vitaz Strait is more complicated. In January, the flow is concentrated between Umboi Islands and New Guinea and flows from the Bismarck Sea into the Solomon Sea, reaching speeds of 110 cm s^{-1} . In June/July, this reverses, but speeds close to 40 cm s^{-1} are still possible (Lindstrom et al. 1990:182). The high ΔR of 333 ± 14 ^{14}C yr for Finschafen at the tip of Huon Peninsula may reflect similar turbulent upwelling (compare with ΔR values of 84 ± 53 and 63 ± 65 ^{14}C yr for U-Th coral from 5 km south of Sialum and -199 ± 50 ^{14}C yr for ^{14}C pre-AD 1955 corals west of Sialum). Support for these observations comes from chlorophyll and sea surface temperature satellite data (Steinberg et al. 2006:10–8), which indicates large-scale turbulence in these channels that varies with wind strength and season. The narrow passage between Umboi Island and the western end of New Britain is considered to play a smaller role in the Solomon/Bismarck Sea interaction (Lindstrom et al. 1990:173). The -199 ± 50 ^{14}C yr value for west Sialum is unusual. This may be a reflection of the date of a post-bomb growth date (AD 1955) and needs retesting. The combined ΔR for the Huon Peninsula is 273 ± 13 ^{14}C yr ($\chi^2_{3:0.05} = 130.6 < 7.81$) with an additional uncertainty of 216 yr reflecting the variability present in this subregion.

Petchey et al. (2005) also hypothesized that upwelling along the New Guinea coastline in response to seasonal reversals in the SEC and NEC may contribute to higher ΔR values at the western edge of the Bismarck Sea. Our results do not inconclusively support this hypothesis. Values for the New

³This value excludes ΔR results from the Solomon Island chain, which is also likely to be affected by wakes and eddies caused by the SEC flow through the islands.

Guinea coastline are fairly low, but variable (subregional $\Delta R = 18 \pm 15$ ^{14}C yr; $\chi^2_{3;0.05} = 18.24 < 7.81$; external variance = 100). Some of the ^{14}C -enriched ΔR values (e.g. -159 ± 46 and -48 ± 74 ^{14}C yr for Muschu Island U-Th dated corals) could be caused by the addition of terrestrial matter via the Sepik and Ramu rivers, offsetting any coastal upwelling. The influence of the Sepik River in particular is far reaching, with a freshwater plume that can stretch to the Admiralty Islands 300 km away (Cresswell 2000; Steinberg et al. 2006:32).

ΔR values collected from islands within the center of the Bismarck Sea (Manus $\Delta R = 18 \pm 13$ ^{14}C yr; Lou Island $\Delta R = 8 \pm 108$ ^{14}C yr; Ninego Group $\Delta R = 59 \pm 13$ ^{14}C yr; Hoskins Peninsula, Kimbe Bay $\Delta R = 45 \pm 19$ ^{14}C yr) (subregional $\Delta R = 40 \pm 8$ ^{14}C yr; $\chi^2_{3;0.05} = 5.15 < 7.81$; external variance = 19) are comparable to values from the northern branch of the SEC (Nauru = 9 ± 5 ^{14}C yr [Guilderson et al. 1998]) and the central South Pacific Gyre average (5.5 ± 21 ^{14}C yr). This value probably reflects calmer, highly mixed surface water conditions at the center of the sea, resulting in a return to typical surface ocean values. Ocean temperatures in this area do, however, indicate that significant upwelling is possible in the lee of Manus Island (Steinberg et al. 2006:33) as the SEC is deflected around New Ireland.

Petchey et al. (2004) have suggested that high values around Kavieng Harbor ($\Delta R = 365 \pm 50$ and 305 ± 110 ^{14}C yr) could reflect equatorial or coastal upwelling along New Ireland. The ΔR of 111 ± 17 ^{14}C yr for New Hanover is more enriched, but in keeping with these observations (subregional $\Delta R = 141 \pm 16$ ^{14}C yr; $\chi^2_{2;0.05} = 25.4 < 5.99$; external variance = 131).

The only ΔR that cannot be explained immediately by oceanographic circulation patterns are from *Tridacna gigas* and *Hyotissa hyotis* recovered from the archaeological site of ECA on Eloaua Island, Mussau (Table 2). Kirch (2001:201–2) calculated an average ΔR correction of -320 ^{14}C yr for shell/charcoal pairs from 2 localities (ECA and ECB), but this value did not give acceptable calibrated ^{14}C ages for *Turbo* sp. gastropods from Mussau. Kirch (2001:202–4) therefore suggested that these surge zone gastropods had been influenced by atmospheric ^{14}C . Petchey et al. (2008) has attributed several negative ΔR values in the South Pacific Gyre to the absorption of atmospheric CO_2 associated with enhanced biological production or wind and wave action in shallow lagoon environments. A similar explanation for surge zone gastropods would, however, result in ΔR values being more enriched in ^{14}C than the bivalves and, therefore, seems to be an unlikely explanation for the observed offset. While we cannot rule out the possibility of sample displacement in the archaeological deposits, this also seems to be an unlikely explanation for the negative ΔR values given the consistency in ΔR offset between the 2 sites (see footnote h to the Tables for ECB ΔR pairs). Given the current available information for Mussau and lagoon environments, we suggest that this unusual ΔR value is an area for further research and, until resolved, remains an unknown error in establishing a reliable shell chronology for the entire Pacific region.

Change over Time

There are relatively few studies of changing R over time in the Pacific. Coral core and foraminifera data from the central Pacific Gyre region suggest negligible change in the surface marine reservoir over the last ~11,000 yr (Bard 1988; Sikes et al. 2000; Paterne et al. 2004) despite several periods of major climatic change noted during this time period. In particular, Haberle and David (2004:166–9) have identified 2 periods of major climate change reflected in burning data and other climate proxies from the Australasian tropics at around 6000–5000 cal yr BP, with increasing El Niño events peaking around 3000–1000 cal yr BP. There is also a growing body of archaeological evidence pointing to the influence of changing climate patterns on human settlement and culture at this time (Allen 2006; Anderson et al. 2006; Nunn 2007; Clark and Reepmeyer 2012). McGregor et al. (2008) sug-

gested that ocean-atmosphere interactions in the western equatorial Pacific triggered El Niño events that pushed equatorial waters eastward, causing marine ^{14}C reservoir variation at Muschu Island between 7220–5850 BP ($\Delta R = -105 \pm 110$ to -265 ± 35 ^{14}C yr; $R = 185 \pm 30$ ^{14}C yr). This was followed by the onset of modern El Niño activity at around 5850–5420 BP that would have enhanced the intrusion of ^{14}C -depleted equatorial waters and a return to modern values. McGregor et al. (2008:220) also identified a shift at around 2065 BP to a ΔR of -135 ± 35 ^{14}C yr ($R = 215 \pm 50$ ^{14}C yr), but data is limited and this value may reflect influence from the Sepik River or recrystallization of the coral. Yu et al. (2010) also noted a shift in ΔR for corals from the South China Sea at around 2800 yr ago, but again, data is limited. In contrast, evidence compiled by Petchey et al. (2009) and Petchey and Clark (2011) indicates stability in ΔR over the last ~3000 yr in the South Pacific Gyre.

For the wider Bismarck Sea/Solomon Sea area, we have identified 4 locations, covering some time depth (AD 1950 [0 BP] to ~3000 yr BP), where $\Delta R/R$ are available; Muschu, Watom Island, and Manus Island (Tables 1–3). Three ΔR values are given from Muschu: -48 ± 74 (at 1940 ± 50 BP); -159 ± 46 (at 2180 ± 50 BP); and 72 ± 60 ^{14}C yr at AD 1912. These 3 values, in an area where river plumes have already been identified as a potential major influence, are outside statistics ($\chi^2_{2,0.05} = 9.43 < 5.99$) and significant variability is also present in the R values ($\chi^2_{2,0.05} = 17.08 < 5.99$). However, deposits dating to ~1650 cal BP ($\Delta R = 321 \pm 103$ ^{14}C yr) and ~3000 cal BP ($\Delta R = 307 \pm 105$ ^{14}C yr) from Watom Island are statistically indistinguishable ($\Delta R: \chi^2_{1,0.05} = 0.01 < 3.84$; $R: \chi^2_{1,0.05} = 0.05 < 3.84$). Similarly, an AD 1954 shell from the Sasi site on Manus Island gave a ΔR of 18 ± 13 ^{14}C yr compared to a value of 8 ± 108 ^{14}C yr in ~2100 cal BP ($\Delta R: \chi^2_{1,0.05} = 0.01 < 3.84$; $R: \chi^2_{1,0.05} = 0.30 < 3.84$). While these limited results are far from conclusive, they support the observation that geographical variation in oceanographic conditions in the Bismarck Sea has more impact on the marine ^{14}C reservoir in the period <3000 yr BP than changing reservoir conditions during this time. This is an area that requires further research.

CONCLUSION

Our results suggest significant variations in the marine ^{14}C reservoir are possible across the Bismarck region. Using ^{14}C from surface ocean carbonates and oceanographic observations, we have identified 6 subregions:

- Region 1 (Coral Sea – including Duke of York Islands, Samarai, Kiriwina Island, and Rabaul), $\Delta R = 38 \pm 14$ ^{14}C yr;
- Region 2 (Huon Peninsula/Vitiaz Strait), $\Delta R = 273 \pm 216$ ^{14}C yr;
- Region 3 (St. George's Channel – west), $\Delta R = 314 \pm 74$ ^{14}C yr;
- Region 4 (eastern New Guinea coastline), $\Delta R = 18 \pm 100$ ^{14}C yr;
- Region 5 (central Bismarck Sea), $\Delta R = 40 \pm 19$ ^{14}C yr; and
- Region 6 (equatorial islands), $\Delta R = 141 \pm 131$ ^{14}C yr.

This suggests that at these locations geographical variation appears to be more extreme than temporal over the last 3000 yr, but more research is needed to confirm this observation and establish the pattern for periods older than 3000 yr BP. Moreover, significant variation between some shell taxa and some near-shore environments is likely and adds uncertainty to these divisions. This is also an area for further research.

Shellfish remains dominate many coastal archaeological sites in the Pacific. Shells therefore provide the greatest opportunity to develop a comprehensive understanding of colonization and changes in settlement patterns and culture since they are sensitive to both the timing and mode of change. Although these new ΔR values go a long way towards improving the reliability of shell dates in this

region, the limited number of ^{14}C dates on charcoal from short-lived materials remains a significant stumbling block to refining both ΔR and chronological issues generally. However, far from being seen as a limitation to obtaining reliable calibrated ^{14}C dates on shell, it should be recognized that the observed complexity within the marine reservoir reflects our growing understanding of ^{14}C in nature, and provides an opportunity for greater dating resolution than has hitherto been available in this region.

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